



## Pre-stack depth migration seismic imaging of the Coral Patch Ridge and adjacent Horseshoe and Seine Abyssal Plains (Gulf of Cadiz): tectonic implications

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**Abstract:** Recently acquired multichannel seismic (MCS) profiles during the SWIM-2006 cruise allow us to characterize the shallow and deep geometry and timing of deformation of the structures comprising the Coral Patch Ridge and adjacent Horseshoe and Seine Abyssal Plains (SAP), in the westernmost Gulf of Cadiz. This region is where the epicentres of the largest instrumental earthquakes occurred, such as the one on 28th February 1969 (Mw 8.0). We present a detailed seismo-stratigraphic and tectonic interpretation of two SWIM-2006 MCS profiles that we have pre-stack depth migrated in order to correct the reflector geometry. Based on drilled wells, we have distinguished six seismo-stratigraphical units (from Triassic to Plio-Quaternary). We have also characterized the 300 km long WNW-ESE lineaments, corresponding to an active dextral strike-slip fault, and the geometry of the Coral Patch Ridge. Finally, present-day active faulting has been observed at the Horseshoe Abyssal Plain and SAP, mainly corresponding to subvertical faults cutting the whole sedimentary sequences up to the surface, some of them associated with earthquake swarms.

**Keywords:** Gulf of Cadiz, multichannel seismics, seismo-stratigraphy, strike-slip fault, thrust fault, seismicity.

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The study area is located in the SW Iberian Margin, in the westernmost part of the Gulf of Cadiz. This region hosts the present-day convergent boundary between Eurasian and African plates (4.5-5.5 mm a<sup>-1</sup>) (Grimison and Chen, 1986; Argus *et al.*, 1989) and is characterized by moderate to intense magnitude seismic activity (Buforn *et al.*, 1995; Baptista *et al.*, 1998; Stich *et al.*, 2005). The Gulf of Cadiz is also the source of the largest seismic events in Western Europe, such as the 1st November 1755 Lisbon Earthquake (Mw 8.5) Johnston, 1996) and the 28th

February 1969 one (Mw 8.0) (Fukao, 1973). Recent estimations of depth and seismic moment tensors (Mw 3.8 to 5.3) for the earthquakes that have occurred in the area show reverse and strike-slip faulting solutions at a depth ranging between 6 and 60 km (Stich *et al.*, 2007).

Numerous marine geophysical cruises have been carried out in the region during the last fifteen years (Sarcori *et al.*, 1994; Banda *et al.*, 1995; Gutscher *et al.*, 2002; Gracia *et al.*, 2003a, b; Terrinha *et al.*,

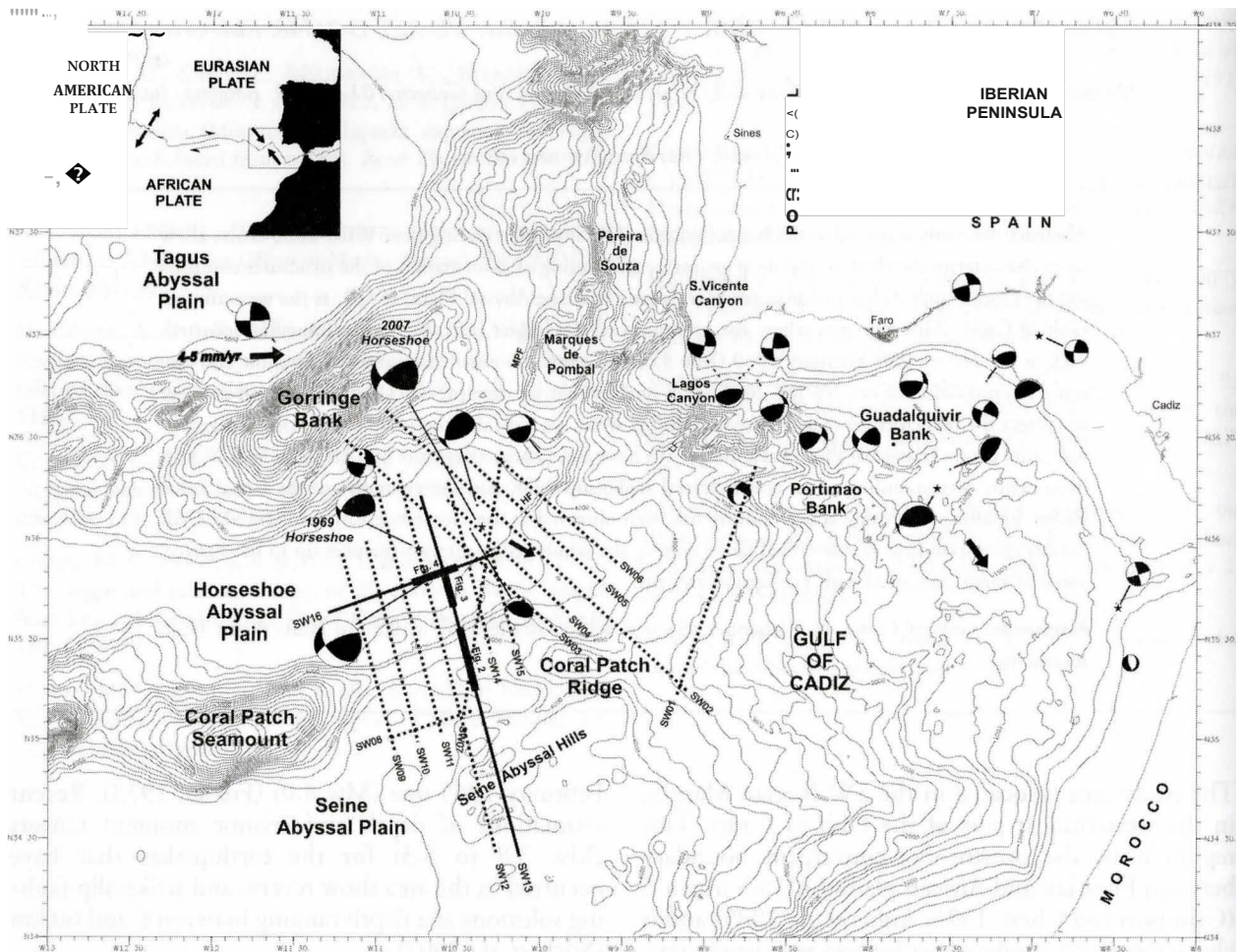
2003; Zicellini *et al.*, 2004); however, there are still some key questions to be solved. For example, what is the nature of the crust and where is the oceanic-continent crust transition located, where is the position of the boundary between the African-Eurasian plates from the Goringe to the Straits of Gibraltar, which are the tectonic sources responsible of the largest earthquakes and tsunamis that have occurred in the area, or what is the real depth-geometry of the active structures identified in the region.

To answer some of these questions we have made a pre-stack depth migration of nine of the sixteen multi-channel seismic profiles acquired during the SWIM-06 marine geophysical cruise. The new seismic images obtained allow us to characterize the real depth and shallow geometry of the active faults and to

quantify the deformation rate and fault parameters. The present study is based on a selection of two profiles (SW13 and SW16) and focuses on: a) a seismo-stratigraphic analysis and calibration from the DSDP site 135, b) the characterization of the upper and lower boundaries of the Horseshoe Gravitational Unit, and c) the obtention of the real geometry of prominent structures, such as the Coral Patch Ridge Fault, the strike-slip SWIM lineaments, and other active faults.

## Data and methods

The SWIM-2006 cruise took place at the external part of the Gulf of Cadiz from 31<sup>st</sup> May to the 14<sup>th</sup> June 2006 onboard the Spanish R/V Hesperides (P. I. E. Gracia). We acquired sixteen high-resolution MCS



**Figure 1.** Bathymetric map of the Gulf of Cadiz with the location of the MCS profiles acquired during the SWIM-2006 cruise (the position of the SW13 and SW16 profiles are indicated in black) and the location of the figures 2, 3, and 4. Plate convergence is shown by black arrows. Fault plane solutions are from Buform *et al.* (1995) and Scich *et al.* (2005, 2007). Inset: tectonic setting of the SW Iberian Margin at the boundary between the Eurasian and African Plates.

profiles (SW01 to SW16) together with Simrad EM120 swath-bathymetry and backscatter, TOPAS sub-bottom profiles, magnetics and gravity data (Fig. 1), covering more than 2700 km of marine geophysical data.

Seismic acquisition was performed using a 10 m array of 8 airguns at 6 m depth producing 1050 c.i. We used an analogical Teledyne streamer with 2.4 km of active section, formed by 96 channels (25 m separation) and cowed at 7 m depth. We recorded the MCS data in SEG-D 48058 rev-1 formats at a sampling rate of 2 ms. The record length was 11 s twtt (two-way travel time) with a shot distance of 37.5 m, with the exception of profile SW01, located in a shallower area, where we obtained a record length of 9 s twtt with a shot interval of 25 m.

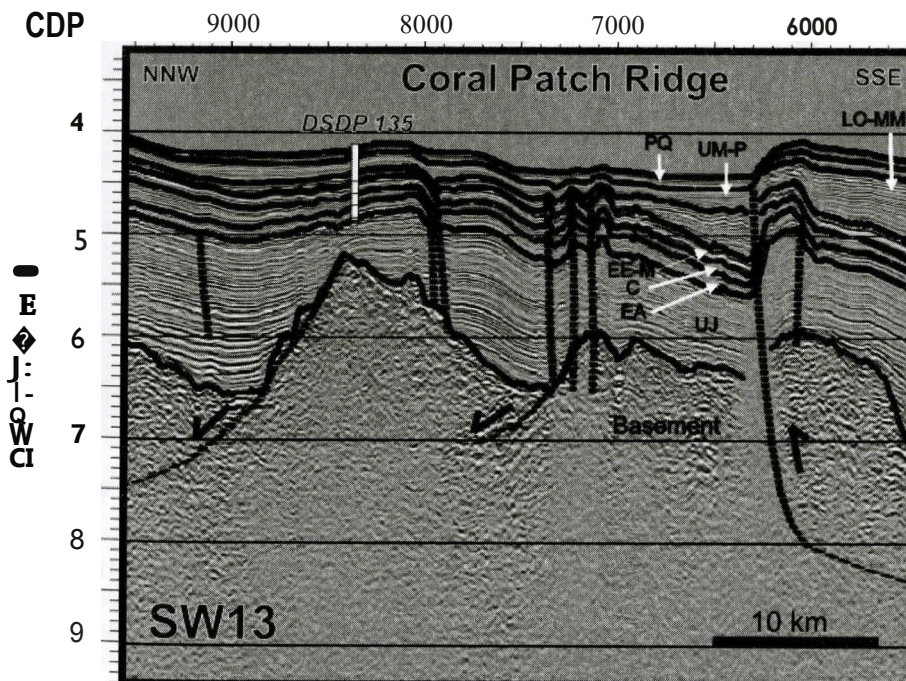
Standard MCS processing was accomplished using PROMAX software, including data resampled from 2 to 4 ms, channel and shot editing, top mutes picked in the shot gather domain, true amplitude recovery, Fx-decon, ensemble predictive deconvolution and geometry CDP gather. A velocity model for Kirchhoff depth migration was performed in 9 MCS profiles (SW01 to SW07, SW13 and SW16) using the SIRIUS software from IFM-GEOMAR (Kiel, Germany) by a depth-focussing error analysis of the MCS data. In the present work we focus on two perpendicular MCS sections: profile SW13, trending NNW-SSE

and 206 km long, and profile SW16, trending WSW-ENE and 103 km long (Fig. 1). The first profile cuts across the Horseshoe Abyssal Plain, Coral Patch Ridge and Seine Abyssal Plain, while the second one cuts along the Horseshoe Abyssal Plain, obliquely to the SWIM Lineament South.

## Results

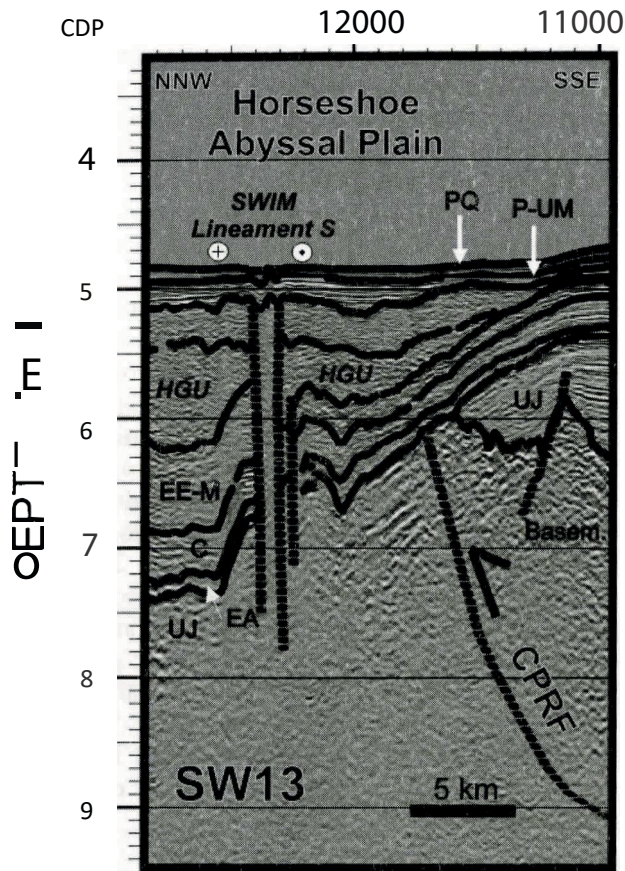
Based on the lithostratigraphic units defined in the 689 m deep well DSDP 135 (Hayes *et al.*, 1972) and calibrated with the depth-migrated seismic profile SW13, which runs exactly on top of this site, we have distinguished six seismo-stratigraphic units in the studied area (Fig. 2): 1) Plio-Quaternary to Miocene Unit (hemipelagites, contourites and turbidite layers), including the Horseshoe Gravitational Unit (Medialdea *et al.*, 2004; Iribarren *et al.*, 2007); 2) Early Eocene-Maestrichtian Unit composed of olive grey to brown silty mudstones, sand layers and brown clays; 3) Cretaceous Unit formed by black and green shales with limestone and chert layers; 4) Early Aptian Unit composed of marls and limestones; 5) Upper Jurassic Unit formed by limestones; and 6) Triassic to Jurassic acoustic basement composed of evaporites and carbonates.

The normal intersection between profiles SW13 and SW16, allows us to obtain a complete perspective of the structures distinguished in the area. For example,



**Figure 2.** Interpreted section of the pre-stack depth migrated MCS profile SW13 between CDPs 5500 and 9500 across the Coral Patch Ridge (1 CDP = 75 m). See location in figure 1. The image shows the location of site

DSDP 135 and the seismo-stratigraphic sequence. Unit I: Late Oligocene to Plio-Quaternary age (LO-PQ) formed by the following sub-units: Late Oligocene to Middle Miocene (LO-MM), Upper Miocene to Pliocene (UM-P), and Plio-Quaternary (PQ); Unit II: Early Eocene to Miocene age (EE-M); Unit III: Cretaceous age (C); Unit IV: Early Aptian age (EA); Unit V: Upper Jurassic age (UJ); and Unit VI: Triassic to Jurassic age, defining the acoustic basement. Vertical exaggeration: 5.



**Figure 3.** Interpretation of the pre-stack depth migrated profile SW13 between CDP's 11000 and 12800 at the Horseshoe Abyssal Plain. See location in figure 1. The image shows the real dip geometry of the Coral Patch Ridge Fault and the SWIM Lineament South. Legend of the seismo-stratigraphic units is included in the caption of figure 2. Vertical exaggeration: 5.

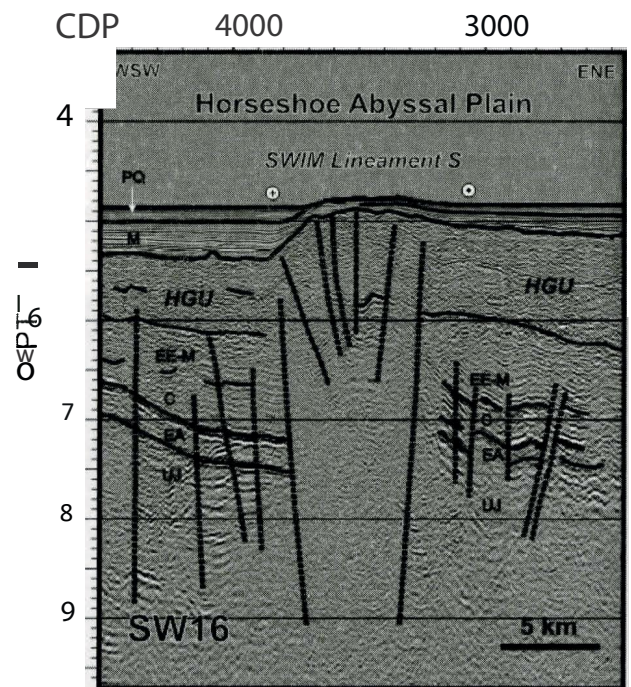
the SWIM Lineament South, corresponding to the bathymetric expression of a WNW-ESE shear zone (Barcolomé *et al.*, 2008; Rosas *et al.*, *accepted*), is crossed by both seismic profiles. SWIM Lineament South is an active strike-slip fault displacing the entire sedimentary sequence up to the surface and locally structured as a flower structure. This lineament runs across the Horseshoe Abyssal Plain to the western part of the Gulf of Cadiz accretionary wedge, reaching down to a depth of 8 km (Figs. 3 and 4). The Coral Patch Ridge (CPR), crossed by profile SW13, is composed of a series of positive reliefs generated by narrow ENE-W SW trending folds and thrusts mainly verging to the NW. An example of these prominent reverse faults is located around CDP 6200 on profile SW13 (Fig. 2), probably reactivated from the Mesozoic rifting phase, causing the vertical uplift and folding of the sedimentary pile above. We have also identified on the MCS profiles that the Jurassic acoustic basement is

structured in half-grabens (Fig. 2) and that the deep geometry of the Coral Patch Ridge Pale corresponds to a large blind thrust (Fig. 3).

The Horseshoe Gravitational Unit (HGU) of Upper Miocene age and filling most of the Horseshoe Abyssal Plain is crossed by both profiles. In profile SW13 (Fig. 3), we observe the lateral boundary between this unit and the CPR, and the resulting wedge geometry. Profile SW16 extends along the Horseshoe Abyssal Plain and allowed us to identify the upper and lower boundaries of the HGU and its internal seismic nature, mainly characterized by chaotic facies (Fig. 4). On both profiles we have observed the deepening of the Mesozoic Units below the HGU (Fig. 3 and 4).

### Discussion and conclusions

The detailed study of the MCS profiles allows us to characterize the deformation sequence of the external part of the Gulf of Cadiz (Fig. 2). We have distinguished between a) pre-rift Triassic and Lower Jurassic



**Figure 4.** Interpretation of the pre-stack depth migrated profile SW16 between CDP's 2500 and 4500 at the Horseshoe Abyssal Plain. See location in figure 1. The section shows the boundaries of the Horseshoe Gravitational Unit (HGU), the Mesozoic units underneath the HGU and the positive flower structure geometry of SWIM Lineament South. Legend of the seismo-stratigraphic units is included in the caption of figure 2. Vertical exaggeration: 5.

deposits composed of evaporites and carbonates (acoustic basement); 6) syn-rift deposits, made up of Jurassic, Early Aptian, Late Cretaceous and Early Tertiary Mesozoic units, mainly composed of terrigenous sediments; c) syn-compressional Miocene-Quaternary deposits divided into Late Oligocene-Middle Miocene, Middle Miocene-Pliocene and Pliocene-Quaternary subunits, composed of hemipelagic, contourites and turbidite layers, and also comprising the Horseshoe Gravitational Unit (HGU) of the Tortonian age (Torelli *et al.*, 1997), a regional marker infilling the Horseshoe Abyssal Plain.

The study of MCS profiles reveals present-day active faulting at the HAP and SAP, mainly subvertical faulting curving the whole sedimentary sequence and the Plio-Quaternary unit showing evidence of recent activity. The MCS SWIM profiles also allowed us to improve our knowledge of the geometry and extension of the Horseshoe Gravitational Unit, and to characterize the sedimentary sequence below it. Finally, we conclude

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chac pre-stack depth migration of MCS profiles is highly recommendable in studies where it is important to know the "real" corrected geometry of the reflectors, such as detailed neotectonic interpretations and calculations of fault seismic parameters, of fundamental importance in order to assess the seismic hazard of the southern Iberian Margins.

## Acknowledgements

The authors acknowledge the support from the European Science Foundation SWIM project (OI-LEG-EMA09F and REN2002-11234E-MAR), Spanish National project EVENT (CGL 2006-12861-C02-02) and the EU-NEAREST project (Ref. 037110). We thank the captain, crew, scientific and UTM technical staff on board the R/V Hesperides, for their assistance throughout the data collection. S.M.-L. and R.B. benefited from a grant of 6 weeks provided by the European Commission transnational access "SALVADORE" project (RITA-CT-2004-505322) to process data at the Seismic Processing Centre of IFM-GEOMAR Kiel (Germany).

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