

Revelation of tin and niobium occurrences in Southern Uis Region of Namibia through a geological reconnaissance study

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Abstract: The paper presents the results of a geological reconnaissance carried out by the author in the Southern Uis region, Namibia. During the present reconnaissance geological studies, close grid sampling was carried out in the investigated area covering about six sq km and a total of 153 individual samples were collected and analysed for their tin, niobium and tantalum contents. The analysis of the samples shows that 14 of them have high positive anomalous values of SnO_2 and 58 samples gave good values of Nb_2O_5 . The samples showing positive values are mainly eluvial in nature and are located in the central portion of the investigated area. As such this area can be mined manually by small group of miners.

Key words: Reconnaissance, tin, niobium, tantalum, Uis, Namibia.

Resumen: Se destacan los resultados de una prospección geológica realizada en una área de 6 km² en el sur de la región de Uis (Namibia). Durante ésta, fueron recogidas un total de 153 muestras y analizadas para determinar su contenido en Estaño, Niobio y Tántalio. Una proporción relativamente importante de las mismas registran valores positivos altos de SnO₂ y buenos de Nb₂O₅ y proceden de depósitos eluviales localizados en la parte central del área estudiada, que podría ser explotada manualmente por pequeños grupos de mineros.

Palabras clave: Prospección geológica, Estaño, Niobio, Tantalio, Uis, Namibia.

Generally tin, tantalum and niobium occur in association with one-another. Pegmatites are supposed to be the largest source of tantalum in the world and an important tin producer. Owing to their petrogenetic significance, the oxides of tin, tantalum and niobium have been extensively studied which throws light on the genesis of granitic pegmatites enriched with rare elements (Cerny et al., 1985; Cerny & Ercit, 1985, 1989; Fuente and Izard, 1998; Fuente et al. 2000; Robles et al., 1999; Bekasmi et al., 2000; Aurisicchio *et al.*, 2001).

Africa is an important resource centre for tin, tantalum and niobium and their mineralization occurs in various countries across the African continent. However, their exploitation is restricted to a few countries. The African countries having a history of tantalum minerals include Namibia, Ethiopia, Mozambique, Uganda, South Africa, and Zimbabwe where alluvial and eluvial deposits are being mined using small-scale techniques.

Based on the regional geological reconnaissance work carried out by the author in an area of about 3000 sq km, some potential areas were identified (Singh, 2007a) and subsequently detailed surface and subsurface exploration works were undertaken for the purpose of geological mapping, surface sampling, trenching, pitting, drilling and bulk sampling (Singh, 2007b, c). Studies carried out by Singh (2007a,b, c) indicate that the pegmatites in this region are LCT type of Cerny's classification given in 1991 but in places they show mixed features between LCT and NYF pegmatites.

The aim of this paper is to present the reconnaissance geological work that has been carried out in about six sq km area of southern Uis region which is a part of Erongo region of Magisterial district Omaruru, Namibia (Fig. 1). This area is included in the Uis-Cape Cross rare element pegmatite belt (GSN, 2000) and has given positive anomalous values in tin (79 ppm of SnO₂), tantalum (0.45 wt. % of Ta₂O₅) and niobium (0.20 wt. % of Nb₂O₅) in previous geological reconnaissances as it is pointed out by Singh (2007a). Taking into account these facts, this study was focussed on the study of pegmatite bodies of this area for assessing the potential occurrences of tin, tantalum and niobium.



Figure 1. Map of Namibia showing Uis, the area under investigation.

Geological Setting

The younger sediments of the Kalahari and Namib deserts cover nearly half of the Republic of Namibia. Nevertheless, the varied geology of this country encompasses rocks spanning from the Archean to the Phanerozoic age and covers more than 2600 Ma of the earth's history. The metamorphic inliers that consist of highly deformed gneisses, amphibolites, meta-sediments and associated intrusive rocks occur in the central and northern parts of the country, and represent some of the oldest rocks of Paleoproterozoic age (2200 to 1800 Ma) in Namibia (Geological Survey of Namibia, 2000).

In the Southern Uis region, meta-sedimentaries, belonging to the Amis River Formation, Swakop Group of 1000-545 Ma (GSN, 2000) crop out. They are the oldest rocks occurring in the area and include quartzite, phyllite, quartz-schist, quartz-mica-schist and mica-schist. They strike in NE-SW to N-S direction and dip towards the south to southeast. Extensive erosion has taken place at regional scale but the resistant rocks can be seen as ridges and hills in the area.

These meta-sedimentaries are intruded by granites and granite gneisses. The intrusive rocks include mainly quartz and alkali-feldspar while tourmaline occurs as an accessory. These intrusive granites and granite-gneisses, normally trend in the strike direction of the meta-sedimentaries. The alkali feldspars in the granite gneiss are medium to coarse grained. At the contact of the intrusive granite with the sedimentaries extensive feldspathization can be observed. The granites, granite-gneisses and meta-sedimentaries have further been intruded by younger pegmatites and aplites, thereby constituting an integral component of the geo-tectonic setting of the area. These pegmatites also trend parallel to the strike of the meta-sedimentaries while in some places they cut across in the form of veins. These veins vary in width from a few metres to sometimes more than 25 metres. However, they do not show noticeable internal zoning. They contain several minerals formed by rare elements in addition to minerals such as quartz, microcline, albite, and muscovite. The accessory minerals include cassiterite, members of columbite-tantalite series of minerals and, in places, zircon and lithium minerals like amblygonite. The rare element pegmatites of the Uis-Cape Cross pegmatite belt were emplaced between 550 and 460 Ma (GSN, 2000). They are considered to be a major source of tin in Namibia. Figure 2 shows the geological map of the investigated area.

Dolerite dykes also occur as intrusive rocks cutting across the formation and intruding all the older rocks. At few places they form ring like structure called 'ring dykes' (Singh, 2007a).



Figure 2. Geological map of southern part of Uis, Namibia.

Method of study

The geological traverses were taken in such a way that the entire area was covered and all the outcrop exposures were studied more specifically. A grid of 100 x 100 m was followed in and around the main pegmatite body of the area which falls in the central portion of the investigated area, while at the peripheral part the samples were collected from the exposed quartz bodies as well as from the eluvials rich in quartz fragments. While collecting the samples within the grid, we took both the exposed rock samples as well as the eluvial samples with quartz fragments. In total 153 samples were collected from the area and were subjected to the analysis for their tin, tantalum and niobium contents. Rock samples were collected from the major exposed quartz-pegmatite bodies and sample locations and were located using a GPS. Figure 3a shows the sample locations of all the 153 samples collected in the present investigation.

The approximate size of the rock samples was about 20 x 20 x 10 cm while about 5 kg of eluvial samples were collected. The collected samples were processed (split, crushed and pulverized) at the 'Analytical Laboratory Service', Windhoek, Namibia and were subsequently analyzed for their tin, tantalum and niobium contents in the Research and Development Centre of NMDC, Ltd., Hyderabad, India, in ICP-MS (Perkin Elmer).



Figure 3. Map showing location of samples in the southern Uis region, Namibia (a) and location of samples (b) with assay values in potential block.

Results and discussions

While carrying out the geological traverses, special attention was focussed on the study of the pegmatite bodies for assessing the potential occurrences of tin, tantalum and niobium. Out of the 153 samples analysed, one gave a value of tin (SnO_2) higher than 3000 ppm, ten between 2000 and 3000 ppm and three between 1000 and 2000 ppm. The rest of the samples showed a tin content of below 1000 ppm. These values reveal their richness due to mineralization. As far as niobium concentra-



Figure 4. Scatter plot of concentrations of Ta₂O₅ and Nb₂O₅.



Figure 5. Scatter plot of concentrations of SnO_2 and $Ta_2O_5 + Nb_2O_3$.

Table I. Details of the samples and their corresponding SnO2, Nb2O5 and Ta2O5 contents. ND-not determinable due to low concentration.

| S.No | Sample | Nature of Sample | SnO2 | Ta2O5 | Nb2O5 | 77 | 77 | Quartzo-feldspathic rock sample | ND | 4 | 114 |
|------|--------|--|-------|-----------|-------|----------|------------|---------------------------------|----|-----|------|
| | No | | | | | 78 | 78 | Quartzo-feldspathic rock sample | ND | 82 | 67 |
| | | | (ppm) | (ppm) | (ppm) | 79 | 79 | Quartzo-feldspathic rock sample | ND | 29 | 133 |
| 1 | 1 | White quartz with black specks | ND | 23 | 37 | 80 | 80 | Quartzo-feldspathic rock sample | ND | 27 | 46 |
| 2 | 2 | Eluvial sample | ND | 23 | 252 | 82 | 82 | Quartzo-feldspathic rock sample | ND | 195 | 62 |
| 3 | 3 | Eluvial sample | ND | 23 | 252 | 83 | 83 | Quartz sample with black specks | ND | 20 | 43 |
| 4 | 4 | Eluvial sample | ND | 23 | 33 | 84 | 84 | Quartz sample with black specks | ND | 24 | 44 |
| 5 | - 4 | Eluvial sample | | 23 | 110 | 85 | 85 | Quartz sample with black specks | ND | 9 | 47 |
| 5 | 5 | Eluvial sample (quartz fragments dominate) | ND | 23 | 119 | 86 | 86 | Quartz sample with black specks | ND | 18 | 39 |
| 6 | 6 | | | 22 | 24 | 88 | 88 | Quartz sample | ND | 20 | 39 |
| 6 | 6 | Eluvial sample | ND | 23 | 34 | 89 | 89 | Quartz sample | ND | 21 | 39 |
| | / | Eluvial sample | 2123 | 23 | 53 | 90 | 90 | Quartzo-feldspathic rock sample | ND | 9 | 176 |
| 8 | 8 | Eluvial sample | 3006 | 23 | 54 | 91 | 91 | Eluvial sample | ND | 21 | 39 |
| 9 | 9 | Eluvial sample (quartz fragmente dominate) | 2344 | 23 | 283 | 92 | 92 | Quartzo-feldspathic rock sample | ND | 21 | 39 |
| | | Eluvial sample (qualiz hagments dominate) | | | | 93 | 93 | Quartz sample | ND | 22 | 39 |
| 10 | 10 | El miel e servele | 2515 | 23 | 34 | 94 | 94 | Quartz sample | ND | 85 | 39 |
| 11 | 11 | Eluvial sample | 2502 | 23 | 102 | 95 | 95 | Quartz sample | ND | 22 | 39 |
| | | Eluvial sample (quartz fragments dominate) | | | | 96 | 96 | Quartz sample | ND | 21 | 39 |
| 12 | 12 | | 2165 | 23 | 209 | 97 | 97 | Quartzo-feldspathic rock sample | ND | 21 | 24 |
| 13 | 13 | Stream bed sediment | 2542 | 23 | 250 | 98 | 98 | Quartz sample | ND | 10 | 39 |
| 14 | 14 | Eluvial sample | 2850 | 23 | 89 | 99 | 100 | Eluvial sample | ND | 10 | 43 |
| 15 | 15 | Eluvial sample | 2227 | 23 | 282 | 100 | 100 | Eluvial sample | ND | 12 | 41 |
| 16 | 16 | Eluvial sample | 2186 | 23 | 74 | 101 | 101 | Eluvial sample | ND | 5 | 289 |
| 17 | 17 | | 2730 | 23 | 142 | 102 | 102 | Quartza foldanathia rack comple | ND | 12 | 102 |
| 18 | 18 | Eluvial sample | ND | 23 | 242 | 103 | 103 | Quartzo-feldspathic rock sample | ND | 20 | 39 |
| 19 | 19 | Eluvial sample | ND | 23 | 253 | 104 | 105 | Quartzo-feldspathic rock sample | ND | 22 | 39 |
| 20 | 20 | Eluvial sample | ND | 23 | 46 | 106 | 106 | Eluvial sample | ND | 24 | 172 |
| 21 | 21 | Eluvial sample | ND | 23 | 217 | 107 | 107 | Quartz sample | ND | 20 | 39 |
| 22 | 22 | Eluvial sample | ND | 23 | 112 | 108 | 108 | Quartz sample | ND | 22 | 39 |
| 23 | 23 | Eluvial sample | ND | ND | 298 | 109 | 109 | Quartz sample | ND | 21 | 39 |
| 24 | 24 | Eluvial sample | ND | ND | 269 | 110 | 110 | Quartzo-feldspathic rock sample | ND | 22 | 36 |
| 25 | 25 | Stream bed sediment | 1526 | ND | 166 | 111 | 111 | Stream bed sediment | ND | 22 | 39 |
| 26 | 26 | Eluvial sample | 580 | ND | 355 | 112 | 112 | Ouartz sample | ND | 20 | 282 |
| 27 | 27 | Eluvial sample(near stream bed) | 1224 | 7 | 146 | 113 | 114 | Quartzo-feldspathic rock sample | ND | 20 | 37 |
| 28 | 28 | Eluvial sample | 1233 | 9 | 146 | 115 | 115 | Quartzo-feldspathic rock sample | ND | 22 | 39 |
| 29 | 29 | Eluvial sample | ND | 22 | 29 | 116 | 116 | Quartzo-feldspathic rock sample | ND | 57 | 39 |
| 30 | 30 | Eluvial sample | ND | | 359 | 117 | 117 | Quartz sample | ND | 18 | 39 |
| 21 | 21 | Eluvial sample | | ND | 200 | 118 | 118 | Stream bed sediment | ND | 12 | 39 |
| 32 | 32 | Stream bed sediment | ND | 82 | 495 | 119 | 119 | Eluvial sample | ND | 17 | 40 |
| 33 | 33 | Eluvial sample | ND | 2 | 193 | 120 | 120 | | | 17 | - 55 |
| 34 | 34 | Eluvial sample(near stream bed) | ND | 2 | 262 | 121 | 121 | Quartzo-feldspathic rock sample | ND | 17 | 39 |
| 35 | 35 | Eluvial sample | ND | ND | 336 | 122 | 122 | Quartz sample | ND | 16 | 39 |
| 36 | 36 | Eluvial sample | ND | ND 442 | 405 | 123 | 123 | Quartz sample | ND | 18 | 39 |
| 37 | 37 | Eluvial sample | ND | ND | 185 | 124 | 124 | Quartz sample | ND | 1/ | 39 |
| 39 | 39 | Eluvial sample | ND | ND | 272 | 125 | 125 | Quartz sample | ND | 16 | 39 |
| 40 | 40 | Eluvial sample | ND | ND | 518 | 127 | 127 | Eluvial sample | ND | 16 | 39 |
| 41 | 41 | Eluvial sample | ND | ND | 582 | 128 | 128 | Quartzo-feldspathic rock sample | ND | 57 | 176 |
| 42 | 42 | Eluvial sample | ND | ND | 718 | 129 | 129 | Quartzo-feldspathic rock sample | ND | 31 | 34 |
| 43 | 43 | Stream bed sediment | ND | ND | 250 | 130 | 130 | Quartzo-feldspathic rock sample | ND | 32 | 33 |
| 44 | 44 | Eluvial sample | | ND | 192 | 131 | 131 | Quartzo-feldspathic rock sample | ND | 27 | 37 |
| 46 | 46 | Eluvial sample | ND | ND | 266 | 132 | 132 | Quartz sample | ND | 26 | 39 |
| 47 | 47 | Eluvial sample | ND | ND | 378 | 133 | 134 | Quartzo-feldspathic rock sample | ND | 24 | 40 |
| 48 | 48 | Eluvial sample | ND | ND | 315 | 135 | 135 | Eluvial sample | ND | 20 | 69 |
| 49 | 49 | Eluvial sample (Near stream bed) | ND | 4 | 139 | 136 | 136 | Quartzo-feldspathic rock sample | ND | 20 | 26 |
| 50 | 50 | Eluvial sample | ND | 2 | 140 | 137 | 137 | Quartzo-feldspathic rock sample | ND | 126 | 33 |
| 52 | 52 | Eluvial sample | ND | ND | 494 | 138 | 138 | Quartzo-feldspathic rock sample | ND | 33 | 29 |
| 53 | 53 | Eluvial sample | ND | ND | 588 | 139 | 139 | Quartz sample | ND | 22 | 27 |
| 54 | 54 | Eluvial sample | ND | 4 | 94 | 140 | 140 | Quartz sample | ND | 23 | 39 |
| 55 | 55 | Eluvial sample | ND | 4 | 202 | 141 | 141 | Quartz sample | ND | 24 | 39 |
| 56 | 56 | Eluvial sample | ND | ND | 316 | 143 | 143 | Quartz sample | ND | 24 | 36 |
| 57 | 5/ | Eluvial sample | | | 192 | 144 | 144 | Quartz sample | ND | 24 | 39 |
| 59 | 59 | Eluvial sample | ND | ND | 448 | 145 | 145 | Eluvial sample | ND | 23 | 136 |
| 60 | 60 | Eluvial sample | ND | ND | 246 | 146 | 146 | Quartzo-feldspathic rock sample | ND | 103 | 36 |
| 61 | 61 | Eluvial sample | ND | ND | 511 | 147 | 147 | Quartzo-feldspathic rock sample | ND | 28 | ND |
| 62 | 62 | Eluvial sample | 244 | 53 | 363 | 148 | 148 | Quartzo-teldspathic rock sample | ND | 18 | 30 |
| 63 | 63 | Eluvial sample | ND | 2 | 139 | 149 | 149 | | | 20 | 34 |
| 65 | 65 | Eluvial sample | 254 | ND | 299 | 151 | 151 | Quartz sample | ND | 63 | 37 |
| 66 | 66 | Eluvial sample | ND | 9 | 116 | 152 | 152 | Eluvial sample | ND | 23 | 47 |
| 67 | 67 | Eluvial sample | ND | 9 | 232 | 153 | 153 | Quartzo-feldspathic rock sample | ND | 22 | 40 |
| 68 | 68 | Eluvial sample | ND | 62 | 207 | | | | | | |
| 69 | 69 | Eluvial sample | ND | 12 | 525 | ND-not o | leterminal | ble due to low concentration | | | |
| 70 | 70 | Eluvial sample | ND | ND | 522 | | | | | | |
| 71 | /1 | Eluvial sample | ND | ND | 339 | | | | | | |
| 73 | 73 | Quartzo-feldspathic rock sample | ND | 12 | 39 | | | | | | |
| 74 | 74 | Eluvial sample | ND | 1 | 209 | | | | | | |
| 75 | 75 | Eluvial sample | ND | ND | 296 | | | | | | |

tion is concerned, 58 samples have shown above 150 ppm of Nb₂O₅. However, the investigated area is almost devoid of tantalum and only one sample has shown above 150 ppm of Ta₂O₅. The details of the analysis showing concentrations of tin, niobium and tantalum is displayed in Table I. Figure 3b provides the assay values of sample along with their locations in the map within the potential block of the present study. From the analytical data it is evident that most of the better values where tin and niobium concentrations are high, are located in the central part of the investigated area. Moreover, the high values are concentrated in the eluvial samples, and the quartz samples from the exposed pegmatite quartz bodies do not reveal anomalous values. The scatter plot (Fig. 4) for concentrations of Ta2O5 and Nb2O5 in this area shows a poor correlation coefficient ($\mathbf{R} = -0.006$). However, the scatter plot (Fig. 5) between SnO₂ versus $Ta_2O_5 + Nb_2O_5$ reveals that with the increase in the concentration of Sn, there is concomitant decrease in the concentration of Nb + Ta and vice versa.

Conclusions

A total of 14 samples out of 153, have given a positive anomalous value of SnO_2 while 58 samples have shown above 150 ppm of Nb_2O_5 . However, the investigated area is devoid of tantalum and only one sample has shown above 150 ppm of Ta_2O_5 . The analysis reveals that both tin and niobium are located in the central part of the investigated area which covers about 0.64 sq km. Considering the mineralization in this part of the region, an area of around $800 \ge 800$ m (~ 0.64 sq km) may be considered ideal for manual quarrying or shallow mining of the eluvials by the small groups of miners in the area, as shown in Figure 3a and b.

Acknowledgements

The author is highly indebted to NMDC (National Mineral Development Corporation, India) for providing him the opportunity for this exploration work in Namibia while working with NMDC. The author extends his gratitude to NIMDC (Nam India Mineral Development Corporation), Windhoek, Namibia for providing the necessary facilities. The author also wishes to put on record the help rendered by High Commissioner of India in Namibia, various Ministries and Departments of Government of Namibia, local administrative and technical authorities of Uis, Namibia.

The author visited the tantalite pegmatite belt of Uis, Namibia twice to accomplish the exploration activity and was accompanied by Shri V.K. Verma, Shri S.N. Siddiquie and Dr. A. Tripathi of NMDC and Dr. Tim Smalley of CAMEC (Central African Mining and Exploration Company), Namibia. Nevertheless, the views expressed in the present paper are exclusively of the author and do not necessarily match with those of NMDC. The author is thankful to Dr. M. Fuertes-Fuente and an anonymous reviewer for critically reviewing the manuscript and to Dr. Josep Poblet, editor of this journal, for giving good suggestions to improve the contents of the manuscript.

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